Katsuhiko Kimura*, Nobuo Ooi** and Shigeru Suzuki***: Evidence of Vegetation Recovery on Yakushima Island after the Major Holocene Eruption at the Kikai Caldera, as Revealed by the Pollen Record of Buried Soils under the Old-Growth *Cryptomeria japonica* Forest

木村勝彦*・大井信夫**・鈴木 茂****: 屋久島のスギ林内で得られた埋没土壌の花粉分析によって 明らかにされた完新世の鬼界カルデラ火砕流噴火後の植生回復

Abstract Pollen analysis was applied to a buried soil underlying the natural *Cryptomeria japonica* forest on Yakushima Island, in southern Japan. This soil developed above the Koya pyroclastic flow, which was deposited by an eruption from the Kikai Caldera, in 6300 y. BP. Pollen assemblages from the buried soil showed three different successional stages of vegetation recovery following the eruption, which were not reported in previous studies. 1: Grassland vegetation, dominated by Gramineae and ferns, indicating early successional stages following damage by pyroclastic flow materials. The age of this vegetation was estimated to be between 5170 and 6300 y. BP. 2: Mixed grassland and forest vegetation with high species diversity, characterized by the presence of *Haloragis* and *Myrica*. 3: Forest vegetation dominated by *Cryptomeria* and *Trochodendron*, corresponding to present surrounding vegetation, but having different species composition.

Key Words: Cryptomeria japonica, Holocene, Koya pyroclastic flow, Pollen analysis, Vegetation recovery

要 旨 屋久島は 6300 年前の鬼界カルデラの噴火で噴出した幸屋火砕流によって植生が大きな影響を受けたとされ ているが,花之江河湿原とその周辺で行われた過去の花粉分析ではそれを示唆するような結果は得られていなかっ た。本研究ではスギ自然林を通る林道脇の露頭に見いだされた火砕流堆積物の上に発達した埋没土壌の花粉分析を行 い,火砕流噴火後の植生の回復過程を示すと考えられる 3 つの異なる花粉組成を得た。即ち、1)イネ科花粉とシダ胞 子が圧倒的に優占する遷移初期の草原植生を示す組成,2)イネ科とシダの優占は変わらないものの,草本,木本とも に種数が多く特に草本ではアリノトウグサが,木本ではヤマモモが目立ち,先駆樹種が定着を始めた状態を示す組成, 3)スギやヤマグルマなど木本花粉が優占し,現在の周辺植生に対応するものの,現在の森林では希な種や,より標高 の低いところに分布する種を多く含む組成である。1)の堆積物の年代は 5170~6300 y. BP の間であると推定され, 噴火後早い時期に埋没した土壌であると考えられた。

キーワード: 花粉分析, 完新世, 幸屋火砕硫, 植生回復, スギ

Introduction

Yakushima Island (Fig. 1) has a famous old-growth natural forest of *Cryptomeria japonica*, including trees more than 3000 years old. The island is known to have been affected by a huge pyroclastic flow that erupted from the Kikai Caldera 50 km northwest of the island about 6300 years ago (Machida & Arai, 1978). The eruption must have destroyed much of the vegetation of the island because thick pyroclastic flow deposits are found everywhere, even around the highest moun-

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Fig. 1 Location of Yakushima Island and Kikai Caldera

tain (1935m a.s.l.) of the island.

There are some palynological studies at and around Hananoego moor (Fig. 2), which developed in a gentle valley at an altitude of 1600m on Yakushima Island (Miyai 1938, Takeoka 1971, Takeoka & Torii 1982, Yasuda 1991). However, all of these data show pollen assemblages very similar to present forest vegetation, that is characterized by the dominance of *Cryptomeria*, and no indications of the eruption have been reported to date.

When examining the pyroclastic flow along a roadside soil section we found there were pockets of buried soil established above the pyroclastic flow deposit. Palynological analysis of the buried soils revealed the existence of vegetation which was not known there before.

Part of this study was published as a project research report for the Japan Environment Agency (Kimura & Suzuki, 1994).

Study site and Methods

Study site

Yakushima island (30°20'N, 130°30'E) is located 100km south of Kyushu, Japan (Fig. 1). Although small



Fig. 2 Topography of Yakushima Island and location of sampling sites. The open circle in the lower figure indicates the location of the ecological study plot of Suzuki & Tsukahara (1987).

(500km² in area), Yakushima is a mountainous island, with a highest peak of 1935m. This mountain, along with the Kuroshio warm current, produces high precipitation on the island. The mean annual rainfall at the Yakushima weather station on the island's east coast is 3820 mm/y, but much more rain is recorded for hillside areas. Annual precipitation in 1982 at the Arakawa dam (670m a.s.l.), near the study site, was 8693mm, while at the weather station it was 4725mm (Eguchi, 1984).

Kikai Caldera, which erupted in 6300 y. BP, lies 50km NNW of the island. This was one of the largest eruptions in the world during the Holocene, and created a huge pyroclastic flow (Koya pfl, Fig. 1) together with a well known and widespread tephra, the Akahoya ash (Machida & Arai, 1978). The pyroclastic flow extended across the sea, reached Yakushima Island, and even deposited at the top of its' highest mountain. The thickest deposit we observed on the island was about 2m. The total volume of ejecta,

Table 1 Vertical distribution of forest vegetation in Yakushima Island expressed by basal area (sum of cross section area of stems). Data modified from Tagawa *et al.* (1984), Suzuki & Tsukahara (1987), Kimura (unpablished), Kohyama *et al.* (1984) and Irikura (1984).

Altit	ude (m)	1500	1200	950	600	300			
Species	Basal Area (m ² /ha)								
Paula coreana Paula coreana Eurya yakushimensis Acer sieboldianum Cornus kousa Rhododendron metternich Clethra barbinervis Cryptomeria japonica Trochodendron aralioides Symplocos myrtacea Pieris japonica Stewartia monadelpha Tsuga sieboldii Abies firma Daphniphyllum macropod Kalopanax pictus Dendropanax trifidus Quercus (Cyclobalanopsis Illicium religiosum Rhododendron tashoi Cletera japonica Camellia sasanqua Eurya japonica Distylium racemosum Symplocos tanakae Symplocos tanakae Symplocos prunifolia Litsea acuminata Persea japonica Ternstroemia gymnanther Pasania edulis Myrsine seguinii Acer morifolium Myrica rubra Ilex rotunda Torreya nucifera Mallotus japonicus Podocarpus nagi Neolitsea acciculata Syzygium buxifolium Castanopsis cuspidata Styrax japonica Daphniphyllum teijsmann Elaeocarpus japonicus	ii lum () acuta () salicina a	$\begin{array}{c} 1.3 \\ 1.0 \\ 0.4 \\ 0.5 \\ 5.2 \\ 0.3 \\ 32.8 \\ 15.1 \\ 1.9 \\ 4.3 \\ 5.6 \end{array}$	0.0 0.7 31.0 12.4 1.0 0.4 20.4 2.6 0.5 3.4 4.2 3.6 5.9 2.1 0.5 0.5 0.2	11.1 7.0 0.4 0.0 1.8 2.1 2.7 0.0 2.2 7.5 5.4 4.6 2.4 1.1 0.4 0.7 12.4 0.7 12.4 0.7 12.4 0.5 0.7 0.3 0.0 0.4	$\begin{array}{c} 3.2\\ 2.0\\ 0.4\\ 3.7\\ 0.2\\ 1.1\\ 0.9\\ 25.0\\ 0.3\\ 2.3\\ 3.7\\ 0.0\\ 0.0\\ 1.5\\ 0.0\\ 7.4\\ 1.8\\ 1.5\\ 0.6\\ 0.5\\ 2.8\\ 0.1\\ \end{array}$	+ 4 2 2 1 6 + + 6 2 1 2 2 2 5 1 1 1			

+ less than 1

including ash, pfl and pumice, is estimated to be over 170km³ (Machida & Arai, 1992).

The present natural vegetation of Yakushima Island is composed mostly of evergreen forests, except for dwarf bamboo scrub at the island's summit. Areas lower than 500m are dominated by *Castanopsis cuspidata* (var. *sieboldii*), those between 500m and 1000m by *Distylium racemosum* and those higher than 1000m by *C. japonica* (Table 1). *Trochodendron aralioides*, which often exhibits a strange epiphytic form on Yakushima, and the two conifers *Tsuga* sieboldii and Abies firma are common in the zone where Cryptomeria dominates. Quercus acuta, Quercus salicina, Illicium religiosum and Rhododendron tashiroi are common at 1200m and below. Vegetation data in Table 1, from 1200m, were obtained from a 4770m² ecological study plot (Suzuki & Tsukahara, 1987) near the sampling sites for this study (Fig. 2).

Samples and Methods

Buried soil samples were collected from roadside cuts at 3 locations, YK-10, YK-100 and YK-101, at



Fig. 3 Geological sections of sampling sites

altitudes of 1050m, 1150m and 1200m respectively (Fig. 2). All are located near a ridge, so that geologically they have not been greatly affected by material from the upper slope. Geological sections of the 3 sites are shown in Fig. 3. At each site, buried soil was found in small concave pockets above the Koya pyroclastic flow. The pyroclastic flow deposit directly covers a weathered granite, which also has a concave surface.

At YK-10, the Koya pfl has a thickness of 20 cm, and black peat and dark brown silty peat, both with thickness of 20 cm, and contain pfl pumice and gravel, fill the depression. The peat is covered by silt with thickness of 30 cm, and top of it is dark brown A horizon soil of present vegetation. Twelve samples were taken in a column from the top soil to the bottom of depression. Eight of these samples were used for pollen analysis. A block of peat was taken from the bottom of the deposit for radiocarbon analysis.

At YK-100, Koya pfl is found only in the bottom of the ditch, with a maximum thickness of 20 cm. Brown silt-sandy silt, of 60 cm thickness, fills the ditch and is covered by light-brown silt with thickness of 40 cm. Samples were taken from the brown silt at 3 points.

At YK-101, the Koya pfl has a maximum thickness of 60 cm, and a black peat is developed on it with a maximum thickness of 25 cm. The peat clearly originates from vegetation which was established on the pfl. The top surface of the peat is clear, but has a rather irregular shape. Gray-brown sandy silt with thickness of 100 cm covers it. Samples were taken from 3 points in the peat layer. A block of peat was taken from the bottom of the layer for radiocarbon dating.

Collected samples were treated with 10% KOH, sieved through a 0.5mm mesh, decanted to remove sand, treated with 50% HF, dehydrated with acetic acid, and treated by the acetolysis method. The residue was saturated in glycerin, and mounted on a slide. All pollen grains and spores on each slide were counted. Frequencies of pollen and spore types were calculated based on total identified pollen and spore counts. Single-grain pollen preparations were preserved for further morphological and taxonomic studies.

Peat samples obtained from YK-10 and YK-101 were radiocarbon dated, using the methanol-LSC method, at the Research Institute for Advanced Science and Technology, Osaka Prefectural University. For the YK-101 sample, acid pretreatment by HCl was used to remove carbonates, possible contaminants from ancient inorganic carbon materials.

site no.	site no. YK-10			YK-100				YK-101						
sample no.	1	- 3	4	5	6	7	9	12	1	2	3	1	2	3
arboreal pollen		-							· · ·			·		
Podocarpus			8	8	14	5	9	1		1	4			
Abies	14	10	6	10	3	7		7	2	16	38	5		
Tsuga	9	3	1	1	4	2	. 1	2	1	2		2		
Pinus			1	4	7	3		1	3	1	1			
Cryptomeria	84	102	37	6	2	5	3	2	362	126	157			
Taxaceae-Cephalotaxaceae									8	6	14			
Salix				2										
Myrica	4	23	58	103	72	45	18	4	15	9	8	4	1	
Carpinus-Ostrya	4	5	2	. 9	2				1	83	54	7		3
Alnus		2	3	6	2	1			7	8	6			
Quercus subgen. Cyclobalanopsis	67	50	59	31	- 39	30	6	8	263	181	149	52	14	22
Quercus subgen. Lepidobalanus					1				2	1	1			
Castanopsis		12	9	5	1	4	3	1	11	3	1	1		
Ulmus-Zelkova		2			2	1			1	1	1		1	
Trochodendron	57	41	37	18	10	8		1	167	103	268	91		1
Illicium	13	7	19	18	3	5			8	9	6		1	1
Corylopsis									1	1	1			
Distylium			4	4	1	3		1	14	2				
cf. Prunus												21		1
Daphniphyllum	34	34	42	17	13	9	7	3	15	22	12	2		
Zanthoxylum			5	6	1	3			2					
Rhus						2								
llex	3								1	1	1			
Acer	4	5	4	13	9	6	1		1		1			
Vitis									1					
Camellia			4	3	3	2					1			
Araliaceae										1	1	2		
Ericaceae	31	118	134	30	13	19			15	36	27	6		
Symplocos	99	121	123	4	4	4		1	169	268	147	16	3	4
Ligustrum		1			1	1								
Viburnum			2	2	3	2								
nonarboreal pollen														
Gramineae	25	63	255	919	590	616	566	566	6	18	28	699	724	627
Cyperaceae	2	2		_		1								
Heioniopsis			3	3										
Chapanadianana Amanada										2	1			
Chenopodiaceae-Amaraninaceae						1			1		1			
Clamatic											1			
Thalictrum			2	28										
Drosara			2											
Rosaceae		1		,	ļ									
Leguminosae		1	0	0	2		1		2	4	2	1		
Haloranis		2	70	201	202	1	10		11	13	9	1	1	1
Crusiferae		0	70	201	203	144	19	1	1	2	2			
Umbelliferae	2	n	1	`	,	2								
Labiatae	5	4	1	2	1									
Compositae subfam Carduoideae		5	20	50	61	20	-							_
Artemisia	6	2	20	20	12	29			I	1	I	2	1	2
Compositae subfam Cichorioideae	0	2	1	12	12	0	2	2	4	2	5	. 3	2	4
fern spores				12	'	2	4	3				17	10	14
Lycopodium	8	80	227	445	173	202	171	154		10	20	15	0	
Osmunda	2	07	227	775	423	505	1/1	134	1	10	20	15	8	1
Trilete spore	2	12	52	52	15	59	20	15	2	4	12	20	-	10
Monolete spore	8	73	108	583	545	326	470	220	10	22	13	202	201	10
	0	15	1.70	565	545	520	472	339	10	23	23	203	390	831
unknown	128	75	102	84	60	54	10	18	87	82	56	32	23	21
total arboreal pollen	423	536	558	300	210	167	48	31	1070	881	800	200	20	27
total nonarboreal pollen	36	83	366	1306	874	804	600	570	26	45	50	723	738	648
total fern spores	18	174	478	1080	983	687	685	514	14	39	56	257	411	842
total identified pollen and spores	477	793	1402	2686	2067	1658	1333	1115	1110	965	1005	1189	1169	1522
total pollen and spores	605	868	1504	2770	2127	1712	1343	1133	1197	1047	1061	1221	1192	1543

Table 2 Counts of pollen and spores for each taxon at each sampling point

Results

Pollen assemblages

Pollen and fern spore samples from three sites, including morphological groups, included 31 arboreal

pollen types, 18 nonarboreal pollen types and 4 fern spore types (Table 2). The vertical distributions of major pollen and spore types are shown in the pollen diagram (Fig. 4). This diagram plots the different



Fig. 4 Pollen diagrams for the three sampling sites, showing major taxa, with proportions higher than 2% for any of the sampling points. Proportions are shown as percentages of the total number of pollen and spores.

species as percentages of total pollen and spores. Moraceae, Rosaceae (excluding cf. *Prunus*) and Leguminosae are considered nonarboreal pollen although they might originate from both trees and herbs. Many of the pollen grains and spores, especially in the top soil of YK-10, were not identified because of bad preservation.

The results show two very contrasting pollen assemblages. Arboreal pollen dominates in YK-100 and the upper samples of YK-10, indicating forest vegetation, whereas, in YK-101 and lower samples of YK-10, nonarboreal pollen and fern spores dominate. The following observations may be made for the three locations sampled.

YK-10: Three local pollen assemblage zones, YK-10-I to YK-10-III are identified. YK-10-I: Samples 12 and 9 show a high occurrence of Gramineae (50.8 to 42.5 %) and monolete spores (30.4 to 35.4). Other pollen and spore types, except *Lycopodium* (ca. 13 %) are rare or absent. YK-10-II: Samples 7 to 4. Arboreal pollen comprises about 10 to 40 % of the total identified pollen and spores. In particular, Myrica (2.7 to 4.1 %) and Quercus subgen. Cyclobalanopsis (1.2 to 4.2 %) exceed 1% for every sample. Gramineae (18.2 to 37.2 %), monolete spores (14.1 to 26.4 %) and Lycopodium (16.2 to 20.5 %) are still common. Arboreal pollen types that characteristically occur in this zone are Podocarpus, Pinus, Distylium, Zanthoxylum, Camellia and Viburnum. Among the nonarboreal pollen types, Haloragis (5.0 to 9.8 %) and Compositae subfam. Carduoideae (1.4 to 3.0 %) are characteristic. YK-10-III: Samples 3 and 1. Arboreal pollen proportions are high, making up 67 % of 10-3, and increasing to 89 % in 10-1. Among arboreal types, Symplocos (15.3 to 20.8 %), Cryptomeria (12.9 to 17.6 %), Q. Subgen. Cyclobalanopsis (6.3 to 14.0 %), Trochodendron (5.2 to 11.9 %), Daphniphyllum (4.3 to 7.1 %) and Ericaceae (14.9 to 6.5 %) are characteristic, and all except Ericaceae increase towards the top of the column. Tsuga (0.4 to 1.9 %) and Abies (1.3 to 2.9 %) have the same tendency although they only increase by a few percent. One fifth of the pollen grains cannot

be identified because of bad preservation.

YK-100: Arboreal pollen comprises more than 90 % of the total pollen and spores. *Cryptomeria* (13.1 to 32.6 %), *Q.* subgen. *Cyclobalanopsis* (14.8 to 23.7 %), *Trochodendron* (10.7 to 26.7 %) and *Symplocos* (14.6 to 27.8 %) are dominant throughout 3 samples. The most dominant types in 100-3, 100-2 and 100-1 are *Trochodendron*, *Symplocos* and *Cryptomeria* respectively. In the lower two samples, *Abies* (0.2 to 3.8 %) and *Carpinus-Ostrya* (0.1 to 8.6 %) are relatively common. *Daphniphyllum* (1.2 to 2.3 %) and Ericaceae (1.4 to 3.7 %) are higher than 1 % throughout the samples. Nonarboreal types are rare, except Gramineae (0.5 to 2.8 %) and monolete spores (0.9 to 2.4 %).

YK-101: This site is characterized by high occurrences of Gramineae pollen and fern spores. Arboreal pollen comprises 17.6 % of sample 101-1, but only 1.7 and 2.1 % in the lower two samples. Among arboreal pollen types, *Q*. subgen. *Cyclobalanopsis* dominates, showing an occurrence of 1 % in lower samples and attaining 4 % in 101-1. In 101-1, *Trochodendron* (7.7 %), cf. *Prunus* (1.8 %), and *Symplocos* (1.3 %) are common trees. Nonarboreal pollen are almost exclusively Gramineae (41.2 to 61.9 %) and monolete spores (17.1 to 54.6 %), followed by some Compositae types such as *Artemisia* (0.2 to 0.3 %) and subfam. Cichorioideae (0.9 to 1.4 %).

Radiocarbon dates

The radiocarbon dates for the peats taken from YK-10 (corresponding to pollen zone YK-10-I) and YK-101 are 3410 ± 80 y. BP (OR-090) and 5171 ± 83 y. BP (OR-123). The former sample was taken from a rather shallow depth (ca. 50 cm) and contains many visible roots, so it would be younger than the actual age of deposition. As the latter sample is from a deeper depth (ca. 120 cm), and acid pre-treatment with HCl has been applied to remove carbonates, the date should be more accurate, but there are still a few rootlets visible in the sample.

Discussion

Successional stages reconstructed from pollen assembleges

Although all the buried soil samples developed above the Koya pyroclastic flow deposit, pollen assemblages obtained from three sites differ, probably because of different ages of deposition. This illustrates the course of vegetation recovery after the eruption.

General vegetation change in this region seems to proceed from a grassland community, composed of Gramineae and ferns (namely, stage 1 : pollen assemblage zone I, YK-10-I & YK-101) through mixed vegetation of grass and trees (stage 2 : pollen assemblage zone II, YK-10-II) to a forest community composed of *Cryptomeria, Trochodendron, Q.* subgen. *Cyclobalanopsis* and *Symplocos* (stage 3 : pollen assemblage zone III, YK-10-III & YK-100). However, it is very difficult to arrange pollen assemblages obtained here on an absolute time scale, as the deposits are not continuous.

Stage 1: The pollen assemblages of YK-10-I and YK-101 indicate that there was a grassland vegetation around the sampling locales. The number of identified taxa from this zone is much smaller than the number from the other zones, and is dominated by Gramineae and ferns. To identify the species of Gramineae, plant opal analysis was carried out on the YK-101 sample after washing it through a 10μ m mesh screen. However, no plant opal was obtained, suggesting that it was not a *Miscanthus sinensis* grassland, the most common volcanic vegetation type in Japan. This grassland might be composed of other Gramineae species which did not produce plant opal larger than 10 μ m.

Among the tree taxa, *Q*. subgen. *Cyclobalanopsis* seemed to establish earlier than others, in both YK-10- I and YK-101. In YK-10-I, *Myrica* and *Daphniphyllum* are also dominant. Since *Q*. subgen. *Cyclobalanopsis* and *Daphniphyllum* are not typical pioneer species, the source of pollen from these trees could be remote forest fragments which survived the eruption.

The radiocarbon ages of YK-10-I and YK-101 are quite different $(3410\pm80 \text{ and } 5171\pm83 \text{ y}. \text{ BP})$ in spite of the similarity in their pollen assemblages. As already mentioned, the former age may to be too young because of modern root contamination. However, the latter date also seems too young, since it is 1100 years after the eruption. In Krakatau Island, Indonesia, dense forests were re-established only 100 years after a catastrophic eruption (Tagawa *et al.* 1985). As the YK-101 sample contains a few roots and carbonates have been removed by HCl pre-treatment, 5171 y. BP is thought to be the minimum age. Thus, the correct age would be between 5171 and 6300 y. BP, but it is difficult to tell when this vegetation was established and how long it lasted.

Stage 2: This zone corresponds to YK-10-II, and is characterized by high diversity but low occurrence (10%) of arboreal pollen taxa. The number of taxa is greater than in zone I, and almost all the taxa occurring in YK-10 are represented in this zone. The vegetation seems to have still been predominantly grassland, but many of the tree taxa characterize this zone, such as *Podocarpus, Pinus, Myrica, Distylium, Zanthoxylum, Acer, Camellia* and *Viburnum*. All of these are minor taxa in the present forest around the site, but some are common in lower elevation forest areas, such as *Distylium* and *Myrica* (see Table 1). Among these, *Pinus, Zanthoxylum* and possibly *Myrica* are pioneer species.

The dominant tree taxa in this zone is *Myrica*, which would be *M. rubra*, a nitrogen fixer. *Myrica* is thought to be a typical pioneer species in volcanic areas in the tropics (Flenley, 1979). In Hawaii, an exotic pioneer *M. faya*, which is also a nitrogen fixer and has seeds dispersed by birds, is reported to expand into areas covered with volcanic ash (Vitousek & Walker, 1989).

Species diversity of nonarboreal taxa is also high and *Haloragis*, *Artemisia* and Compositae subfam. Carduoideae are typical in this zone. *Haloragis micrantha*, the only possible species of the genus, which is a small herb growing on wet but open places, might have grown in the small ditch where the buried soil was formed. The local vegetation would have been an open grassland, composed of many herbaceous species, with many tree species, including pioneers, invading the surrounding area. *Cryptomeria* was still a minor species even among tree taxa.

One of the problems associated with this kind of buried soil sample is the possibility of mixing of two or more different pollen containing soils that originated in completely different times. This could produce a fake flora. However, in YK-10 some taxa, such as *Haloragis*, Compositae subfam. Carduoideae and *Myrica*, characterize zone II and decrease in both the upper and lower zones, suggesting that the assemblage of pollen and spores in this zone is not a mixture of soils from zones I and III, but reflects the actual flora that existed at a specific time in the past.

Stage 3: High occurrence of arboreal pollen in zone III clearly indicates that the vegetation has already changed to forest. This zone includes samples YK-10-III and YK-100. YK-10-III is the top soil from the present forest, but YK-100 is buried 40-100 cm and thus might have been protected from the effects of sub soil animals. Pollen assemblages from this zone correspond to the present vegetation around the study area, but there are three interesting discrepancies between them.

1. Some of the pollen taxa found in this zone are very rare, or even not recorded in the present flora of Yakushima. Carpinus-Ostrya is relatively abundant in YK-100, but it is very rare in the present flora (only C. laxiflora is recorded for this taxa in the flora list of Mitsuta & Nagamasu (1984)). Alnus (only A. firma is in the list) and Quercus subgen. Lepidobalanus (Q. acutissima) are found in this zone, although not frequently, are also rare in the present flora. Ulmus-Zelkova is not recorded in the present flora list. Takeoka & Torii (1982) detected pollen of Zelkova. It is interesting that some of the species that survived the catastrophic eruption and expanded after it, were extinguished or almost extinguished in the process of vegetation recovery. This implies that species extinction does not always occur due to the catastrophic event itself.

2. Many of the taxa in the pollen flora are distributed in lower elevation areas in the present forest vegetation. Q. subgen Cyclobalanopsis, one of the dominant taxa in YK-100, is distributed mainly below 1000m a.s.l. and is not a dominant species in the forest around YK-100 (Table 1). The pollen taxa Myrica, Castanopsis- Pasania, and Podocarpus would correspond to the species Myrica rubra, C. cuspidata / P. edlus and P. nagi /P. macrophyllus. Their present vertical distribution is centered in elevations below 500 m and usually never goes higher than 1000m. A similar phenomenon is reported for pollen records from the Hananoego moor (Fig. 2). Takeoka (1971) ascribed this to wind transport from lower elevation areas, but it is possible that these species were actually distributed around the site.

3. The low occurrence of *Tsuga* and *Abies*. As *T*. sieboldii and A. firma are dominant tree species in the present forest near the sampling points (Table 1), the occurrences of the two species in the pollen record appear to be rather low. The highest occurrence of Abies, found at the bottom of YK-100, is 3.8%, and that of Tsuga, in the top soil of YK-10, is only 1.5%. Takeoka & Torii (1982) reported a significant increase of Tsuga in the surface sample. Expansion of Tsuga and Abies seems to have taken place rather recently. In general, *Tsuga* more often grows on ridges with shallow soil than on slopes with deep soil. So, expansion of Tsuga might not have started until the pyroclastic flow deposit was eroded and bare rocks emerged on the surface. Extensive cutting of Cryptomeria trees during the Edo era (100-350 years ago) might have accelerated this erosion process.

Comparison with previous studies

The pollen assemblages of Zones I and II in this study are completely different than previously studied pollen assemblages from the *Sphagnum* peat bog in Hananoego moor (Miyai, 1938; Takeoka, 1971; Takeoka & Torii, 1982; Yasuda, 1991). The earlier data are much like the pollen assemblage of zone III, which shows forest vegetation. These differences may result from different sample ages rather than different sampling locations.

As the Hananoego moor is located in a gentle valley, pyroclastic flows might have deposited much deeper than in other areas. Erosion would likely have started soon after deposition and continued down to the surface of the bed rock, because no thick pfl deposit was observed under the bog (Sohma, 1984: Fig. 5). Radiocarbon ages of peat samples obtained from the bottom of Hananoego moor are 3380 y. BP (Takeoka & Torii, 1982), 2600 y. BP (Sohma, 1984) and 2460 y. BP (Yasuda, 1991) indicating that initiation of the deposition was rather late. Sohma (1984) ascribed this late start of deposition to unstable slope conditions caused by heavy rainfalls which started around 2000 y. BP.

On ridges and upper slopes erosion would have been stopped by the relatively early recovery of the surrounding vegetation, and some soils of this age, especially those in gully like microtopography, might have been buried by occasional small scale slope failures. This would be the origin of buried soils in YK-10 and YK-101. This kind of buried soil is useful because it is commonly found in other volcanic areas.

Features of soil pollen analysis

In the case of pollen analysis of soils, as opposed to *Sphagnum* peat bog samples, surface soil on which pollen falls is not a newly produced pollen free material, but an aged material that always contains pollen from former vegetation, except when it is in very early successional stages. So, pollen assemblages of upper soils would be more or less of a mixture of different ages. Also, vertical transport by soil animals, or leaching and preservation of pollen, would be important considerations in the soil analysis.

We haven't discussed the preservation of pollen in the soil, however, great differences in preservation among species are known. Havinga (1984) studied deterioration of pollen and fern spores in different fill materials and showed differences in preservation rate among species and fill materials. In the extreme case, all the *Lycopodium* spores survived in river clay for 10 years, whereas 99 % of *Myrica* deteriorated within a few years. The differences were greater in river clay and leaf mould than in peat. Thus the actual dominance of *Myrica* in YK-10-II might have been much greater than we observed.

In spite of these problems, however, buried soil provides an advantage in pollen analysis, because peat bogs develop in very restricted locations, especially in steep mountain areas as on Yakushima.

The effects of erruption to the vegetation

This study found that vegetation around the study sites became grassland some time after the pyroclastic eruption in Kikai Caldera. However, some portion of forest must have survived the catastrophe since Yakushima has a rich flora including many endemic plant species. Mitsuta & Nagamasu (1984) noted that on Yakushima, there were many primitive plant species which could only survive in stable natural forest conditions.

Effects of the eruption still exist in the present forest vegetation because there are extensive pyro-

clastic flow deposit remnants on the island. These are important as a source of mineral nutrients for plants and as mother material for the present forest soil. However, they are being gradually eroded by rainfall, which will make the species composition of the forest change in the future.

Acknowledgments

We thank Dr. Setsuko Shibata of the Research Institute for Advanced Science and Technology, Osaka Prefectural University for her radiocarbon measurements of our peat samples. We also thank Donald M. Thieme, Rodger Sparks and Nicola Redvers-Newton, members of the Radiocarbon Mailing List of Arizona Univ., for their replies to my question about contamination problems in soil carbon dating. This study was partly supported by the project of Japan Environment Agency "Long term ecological studies in the Yakushima Wilderness Area and its surrounding areas, 1993-1994".

References

- Eguchi, T. 1984. Climate of Yaku-shima Island, especially regionality of precipitation distribution. "Research report of Yaku-shima Wilderness Area" (ed. Nature Conservation Bureau, Japan Environment Agency), 3-26 (in Japanese).
- Flenley, J.R. 1979. The equatorial rainforest: a geological history. 162pp., Butterworth, London.
- Havinga, A.J. 1984. A 20-year experimental investigation into the differential corrosion susceptibility of pollen and spores in various soil types. Pollen et Spores, 26: 541-558.
- Irikura, S. 1984. Altitudinal zonation of vegetation at west part of Yaku-shima Island. "Research report of Yaku-shima Wilderness Area" (ed. Nature Conservation Bureau, Japan Environment Agency), 353-374 (in Japanese).
- Khoyama, T., Sakamoto, K., Kobayashi, T. & Watanabe, R. 1984. Structure of tree community of a primary lucidophyll forest in the Koyoji basin. "Research report of Yaku-shima Wilderness Area" (ed. Nature Conservation Bureau, Japan Environment Agency), 375-397 (in Japanese).

Kimura, K. & Suzuki, S. 1994. Pollen analysis of

buried soil under natural *Cryptomeria* forest in Yakushima Island and effects of pyroclastic-flow eruption of Kikai Caldera to the vegetation. "Long term ecological studies in the Yakushima Wilderness Area and its surrounding areas" (eds. Nature Conservation Bureau of Environment Agency and Nature Conservation Society of Japan), 169-177 (in Japanese).

- Machida, H. & Arai, F. 1978. Akahoya Ash-A Holocene widespread tephra erupted from the Kikai Caldera, south Kyushu, Japan. The Quat. Res. (Tokyo), 17: 143-163 (in Japanese).
- Machida, H. & Arai, F. 1992. Atlas of Tephra in and around Japan. 276pp. Univ. of Tokyo Press, Tokyo (in Japanese).
- Mitsuta, S. & Nagamasu, H. 1984. Flora of vascular plants (ferns, fern allies and phanerogams) of the Yaku-shima Wilderness Area. "Research report of Yaku-shima Wilderness Area" (ed. Nature Conservation Bureau, Japan Environment Agency), 103-286 (in Japanese).
- Miyai, K. 1938. Pollen analytical studies on the moor of Yakujima. J. Jpn. For. Soci., 20: 400-410 (in Japanese).
- Sohma, H. 1984. Changes in stability of mountain slopes of Yakushima Island analyzed by variation of inorganic content in peat deposits. J. Geogr., 93: 25-35 (in Japanese).
- Suzuki, E. & Tsukahara, J. 1987. Age structure and regeneration of old growth *Cryptomeria japonica* forests on Yakushima Island. Bot. Mag. Tokyo, 100: 223-241.
- Tagawa, H., Suzuki, E., Fuji, A., Fujii, K., Ohira, H., Susukida, J. & Shioya, K. 1984. Altitudinal change and regeneration structure of a natural *Cryptomeria japonica* forest. "Research report of Yaku-shima Wilderness Area" (ed. Nature Conservation Bureau, Japan Environment Agency), 481-500 (in Japanese).
- Tagawa, H., Suzuki, E., Partomihardjo, T. & Suriadarma, A. 1985. Vegetation and succession on the Krakatau Island, Indonesia. Vegetation, 60: 131-145.
- Takeoka, M. 1971. Studies on the distribution of the natural stand of SUGI (*Cryptomeria japonica* D. Don) in Kyushu, Japan (III). Bull. Kyoto Prefec-

tural Univ. For., 16: 29-33 (in Japanese).

- Takeoka, M & Torii, A. 1982. Changes of the forests around the Hananoego moor at Yakushima, Kyushu, based on pollen analysis. 日林論, No. 93, 283-284 (in Japanese).
- Vitousek, P. M. & Walker, L. R. 1989. Biological invasion by *Myrica faya* in Hawai'i: plant demog-

書評(新刊紹介):地球環境問題への国際的取り組みの特集

今年3月に社団法人環境情報科学センターから,環境 情報科学の25巻1号として「環境の四半世紀-地球環 境-」と題した特集号が刊行された。ひじょうに多岐に わたる地球環境問題への国際的な取り組み方およびその 歴史が概括的,かつ今日的問題をもふんだんに含むかた ちで特集されている。総ページ150のうち113が特集記 事に当てられており,量的にも圧巻である。この特集は 「環境の四半世紀」と銘打った通年特集企画の第1弾「地 球環境編」であり,続いて「環境のデザイン編」が5月 に刊行される予定である。

この特集は、編集委員の袴田共之の巻頭言にあるよう に「環境問題,とくに地球環境問題のむずかしさのひと つは、そこに包合される問題群の個々の現象の認識から 政策実現までの全段階において、広範でかつ深い科学的 知見を要求されるという点にある。(中略)わが国民各階 層の関心を上滑りの、いわゆるブームに終わらせてしま うのでなく、問題解決のうねりとして実体化していくた めには、広範で深い科学的知見を全体として整理して把 握しておく必要がある」という認識を背景として企画さ れた。前半では、この四半世紀におけるおもな流れと現 時点の主要な側面をそれぞれの分野の専門家が概説し、

その構成は、地球環境の四半世紀(橋本),わが国の環境 政策(森島),地球環境問題-地球科学と国際政治との融 合(米本),気候変動に関する政府間のパネルの活動と今 後(西岡),国連気候変動枠組条約の制定経緯と取り組み の現状(松村),地球環境問題の意識の変遷(石),南北 問題からみた地球環境-グローバル化における環境と貧 困(勝俣),NGOの活動と地球環境問題-「マングロー ブ植林行動計画」のベトナム・プロジェクトを事例とし て(向後)の8論説と,「地球環境はいま」と題する7論 説すなわち,地球温暖化研究の過去と現在(松野),オゾ ン層の破壊(富永),酸性雨-これまでとこれから(原), 熱帯林の消失と劣化はいつまで続くか(熊崎),砂漠化-新たな対応への期待(門村),生物多様性のもつ意味(安 野),海洋汚染(宮崎)からなる。

後半の構成は、日本学術会議の地球環境研究へのかか

raphy, nitrogen fixation, ecosystem effects. Ecological Monographs, 59, 247-265.

Yasuda, Y. 1991. Influence of the vast eruption of Kikai Caldera volcano in the Holocene vegetation history of Yakushima, south Kyushu, Japan. Japan Review, 2: 145-160.

(Accepted on November 22, 1995)

わり-第14期・第15期の地球環境関連特別委員会の活 動(吉野),第16期の地球環境と人間活動特別委員会の 活動状況(平田)、地球環境に関する研究体制(北村)の 3つの日本学術会議における地球環境研究の概要,学術 審議会建議「地球環境科学の推進について」(平川),地 球環境研究遂行の見取り図(西岡)の2つの概説,およ び3つの大きなプログラムである IGBP(地球圏-生物 圏国際協同研究計画), HDP (地球環境変化の人間次元の 研究計画),WCRP (気候変動国際協同研究計画)の個々 のプログラムの国際動向に関する次のような25の概説 からなっている。IGBP では、IGBP の国際動向、日本に おける IGBP 研究の概要, IGAC, JGOFS, LOICZ, GCTE, BAHC, GAIM, IGBP-DIS, PAGES, LUCC, HDP では、HDP の国際動向、地球環境問題の「人間的 次元」とは何であったか、DEPS, HDP-GIS, ED, EMA, WCRP では, WCRP の国際動向, 日本における WCRP, TOGA, WOCE, GEWEX-CLIVAR, GEWEX-GAME, ACSYS, SPARC が取り上げられている。これらに START (解析・研究・研修システム)と GCOS (全地球 気候観測システム)の2つの大きなプロジェクトの概説 が加えられている。最後に年表地球環境の四半世紀が8 ページにわたってまとめられている。

どの論説・概説をとってみても簡潔でコンパクトにま とめられているので、興味のある項目について調べるに も便利であるし、研究体制の構造や各プロジェクトの関 わり方、これまでの経緯が分かり易い図・表にまとめら れているので理解し易い。

地球環境問題は地球社会が抱える深刻な問題である。 これまでの取り組みと現状を理解し、これからのわれわ れの関わり方を考えるのに、この特集は大いに役立つも のと思われる。

なお,この特集は一部 2575 円で購入することができる。環境情報科学センターは☎ 03-3265-3916, FAX 03-3234-4307, 振替 00190-3-75900 である。

(辻 誠一郎)